Article

Evolution

Evolution of stomatal and trichome density of the *Quercus* delavayi complex since the late Miocene

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Abstract A fossil oak species, *Quercus tenuipilosa* Q. Hu et Z.K. Zhou, is reported from the upper Pliocene Ciying Formation in Kunming, Yunnan Province, southwestern China. The establishment of this species is based on detailed morphologic and cuticular investigations. The fossil leaves are elliptic, with serrate margins on the apical half. The primary venation is pinnate, and the major secondary venation is craspedodromous. The tertiary veins are opposite or alternate-opposite percurrent with two branches. The stomata are anomocytic, occurring only on the abaxial epidermis. The trichome bases are unicellular or multicellular. The new fossil species shows the closest affinity with the

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extant O. delavavi and the late Miocene O. praedelavavi Y.W. Xing et Z.K. Zhou from the Xiaolongtan Formation of the Yunnan Province. All three species share similar leaf morphology, but differ with respect to trichome base and stomatal densities. Q. tenuipilosa, Q. praedelavayi, and Q. delavayi can be considered to constitute the Q. delavayi complex. Since the late Miocene, a gradual reduction in trichome base density has occurred in this complex. This trend is the opposite of that of precipitation, indicating that increased trichome density is not an adaptation to dry environments. The stomatal density (SD) of the Q. delavayi complex was the highest during the late Miocene, declined in the late Pliocene, and then increased during the present epoch. These values show an inverse relationship with atmospheric CO_2 concentrations, suggesting that the SD of the Q. delavayi complex may be a useful proxy for reconstruction of paleo-CO₂ concentrations.

Keywords Quercus delavayi complex \cdot Quercus tenuipilosa \cdot Morphology evolution \cdot Neogene \cdot CO₂ concentration

Plant morphology is determined by the interactions between genetic and environmental factors. Understanding the impact of environmental change on plant traits is an important issue in evolutionary biology. As the only direct evidence of past life, fossils provide important information on the interactions between plants and environmental change. For instance, "xeromorphic" cuticular features such as epidermal trichomes, sunken stomata, and stomatal furrows are commonly considered as adaptations to aridity, increasing the boundary layer resistance and consequently limiting transpiration [1, 2]. By comparing the traits of closely related fossil species from different geological

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periods, we can better understand the responses of plant traits to environmental change [3-6].

Growing plants in artificially controlled environments is the most common method to study the responses of plant traits to environmental change. By this method, Beerling and Woodward found that the stomatal densities and indices of Hedera helix L. and Chlorophytum comosum (Thunb.) Baker decreased with the doubling of the CO₂ concentration [7]. Another method is to compare traits of the same species from different climatic conditions instead of controlling the environmental factors artificially. By investigating the stomatal density (SD) of Nothofagus soladri (Hook. f.) Oerst., Kouwenberg et al. [8] found a positive relationship between SD and elevation along altitudinal gradients. As the concentration of CO₂ decreases when the elevation increases, the SD of N. soladri is negatively correlated with CO₂ concentration. On the other hand, Zhou et al. [9] and Hu and Zhou [10] indicated that the SD of Quercus pannosa Hand.-Mazz. is positively associated with CO₂ concentration. Another widely used method to investigate the response of plant morphology to climatic change across a large scale is to compare closely related fossils from different geological periods. Franks and Beerling [11] studied the relationship between the stomatal size, intensity, and inductance of plants belonging to the same group from different time periods in relationship to the atmospheric CO_2 concentration. The results showed that the atmospheric CO₂ concentration is the main factor driving the evolution of stomatal conductance [11]. This method gives direct evidence as to how plant traits correspond with climatic change. However, few studies have been carried out, as most groups lack such a complete fossil record.

Quercus subgen. *Cyclobalanopsis* Oerst. (Fagaceae) are evergreen trees mainly occurring in subtropical to tropical regions in East and South Eastern Asia [12–15]. China is a major center of diversity for subgen. *Cyclobalanopsis* and contains 69 species, which represents 62 % of the total species from all over the world, and are often found as dominant elements in subtropical to tropical broad-leaved evergreen forests [14]. *Q.* subgen. *Cyclobalanopsis* has an abundant fossil record in East Asia during the Cenozoic [16–19]. In Yunnan Province, *Cyclobalanopsis* fossil records are also rich starting in the Oligocene period, including impressions, compressions, and fossil fruits [18–20]. These fossils are comparable with extant species making them suitable to study the evolution of traits in different geological periods.

In this study, we describe a new *Quercus* species from the Pliocene Ciying Formation based on well preserved fossil leaves. The fossil species is comparable with extant *Q. delavayi* and *Q. praedelavayi* Y.W. Xing et Z.K. Zhou from the Late Miocene Xiaolongtan Formation in gross morphology, but differs from these two species in stomatal and trichome densities. We consider these three species as the *Q. delavayi* complex. Based on detailed comparisons of the SD and the trichome density of the *Q. delavayi* complex, we discuss the evolution of the micromorphology of the *Q. delavayi* complex in relation to the changing environment.

1 Materials and methods

1.1 Materials

The fossils were collected from the Pliocene Ciying Formation exposed at Hunshuitang county, located ca. 25 km northeast of Kunming, the capital city of Yunnan Province, southwestern China (Fig. 1). The Ciying Formation can be divided into two parts. The lower layer is composed of gray sandstone, and the upper layer is composed of silty mudstone and sandstone [21]. Current fossils were found in the upper member of the Ciying Formation, which also yields abundant fossil plants and shells. Based on the lithological sequences, the Ciying Formation was considered as belonging to the late Pliocene age [22–24].

Specimens of extant *Quercus delavayi* for comparison are from Herbarium of Kunming Institute of Botany, Chinese Academy of Sciences KUN. The specimens used for epidermis analysis were collected by M. Deng and J.J. Hu in 2010 (Table 1).

The descriptions of Xing et al. [19] of the late Miocene *Quercus praedelavayi* fossils from the Xiaolongtan Formation of Xundian, Yunnan were used. Information on the studied specimens is gathered in Table 1.



Fig. 1 Map showing the fossil locality (leaf) of Quercus tenuipilosa

Voucher no.	Locality	Age	Coordinates	Altitude (m)	Year of collection
HLT450	Kunming	Late Miocene	N25°25′, E102°51′	2,200	2007
HST856	Kunming	Late Pliocene	N25°06', E102°57'	2,102	2010
DH008	Jingdong	Extant	N24°26', E100°54'	1,431	2010
DH020	Binchuan	Extant	N25°54', E100°25'	1,862	2010
DH025	Binchuan	Extant	N25°57', E100°22'	2,630	2010
DH029	Binchuan	Extant	N25°57', E100°22'	2,480	2010
DH030	Binchuan	Extant	N25°56', E100°24'	2,050	2010
DH032	Eryuan	Extant	N26°19′, E099°59′	2,300	2010
DH034	Jianchuan	Extant	N26°22′, E099°58′	2,541	2010
DH037	Lijiang	Extant	N26°38', E099°57'	2,274	2010
DH044	Lijiang	Extant	N26°53', E100°14'	2,442	2010
DH067	Lijiang	Extant	N26°52', E100°01'	2,159	2010
DH074	Eryuan	Extant	N25°56', E099°49'	1,932	2010
DH075	Eryuan	Extant	N25°54′, E099°49′	1,763	2010
DH076	Yangbi	Extant	N25°50', E099°53'	1,689	2010
KUN0094673	Jinggu	Extant	?	1,600	?
KUN0396795	Zhongdian	Extant	?	1,900	1981
KUN0396899	Weixi	Extant	?	1,920	1981
KUN0449347	Fumin	Extant	?	1,670	?
KUN0449364	Jiangchuan	Extant	?	1,950	1989
KUN0449366	Heqing	Extant	?	1,900	?
KUN0449389	Songming	Extant	?	1,920	1956
KUN0449391	Shuangbai	Extant	?	2,090	?
KUN0449397	Kunming	Extant	?	1,035	1942
KUN0468641	Lufeng	Extant	?	1,900	1982
KUN0504244	Yanshan	Extant	?	1,200	?

Table 1 The voucher information of fossils and extant species

? means unknown

Table 2 Comparison of the leaf macro- and micro- morphology of Quercus tenuipilosa with its closely related species

Species and references	Age	Gross morphology			Micro-morphology				
		L:W	Ν	Р	S	Т	U&M	<i>TD</i> (n mm ⁻²)/ <i>TI</i> (%)	<i>SD</i> (n mm ⁻²)/ <i>SI</i> (%)
Quercus praedelavayi [19]	Late Miocene	3	13–15	+	+	+	~	282/?	1,139/?
Q. aff. delavayi [37]	Late Miocene	4	10	\sim	~	+	+	192/5.2	465/11.7
Q. tenuipilosa	Late Pliocene	2.5	13	+	+	+	+	250/10.2	672/19.1
Q. delavayi [14]	Extant	2.8–4	10–14	+	+	+	+	194/3.6	931/15.2

L leaf length, W leaf width, N number of pairs of secondaries, P straight primary vein, S secondary vein alternate, T tertiary vein branch, U&M unicellular and multicellular trichome bases on the adaxial epidermis, TD trichome density on the abaxial epidermis, TI trichome index on the abaxial epidermis, SD stomatal density on the abaxial epidermis, SI stomatal index on the abaxial epidermis, + yes, ~ no, ? unknown

1.2 Methods

- (1) *Epidermis characteristics of the fossil species* The procedure of making fossil epidermis slides followed the methods of Ye [25], Leng [26], and Ma et al. [27].
- (2) *Epidermis characteristics of the extant Quercus delavayi* We followed the method of Stace to make the epidermis slides of extant *Q. delavayi* [3].
- (3) *Making cleared leaves* For the examination of the leaf morphology of extant subgen. *Cyclobalanopsis* leaves, we followed the methods of Hickey and Wolfe [28] to make cleared leaves (Table 2).
- (4) *Trichome density* Instead of counting trichomes directly, the trichome bases were counted. Three visual fields under $20 \times$ magnification were used to calculate the trichome density (Table 2).



 Table 3
 The taxon-characteristic matrix for the PCoA analysis

Characters	Species						
	Quercus praedelavayi [19]	Q. aff. delavayi [37]	Q. tenuipilosa	Extant Q. delavayi [14]			
Lanceolate leaf	0	1	0	0			
Elliptic leaf	1	0	1	1			
Leaf length ≥ 10 cm	1	0	0	1			
Leaf width $\geq 2 \text{ cm}$	1	0	1	1			
Length:width <5	1	0	1	1			
Number of secondary pairs >10	1	0	1	1			
Primary vein straight	1	0	1	1			
Secondaries alternate and opposite	1	0	1	1			
Trichomes on adaxial epidermis	0	1	1	1			
Anomocytic stomata	1	1	1	1			
Uni- and multi- cellular trichome bases	0	0	1	1			
Trichome 0 density <200 n mm ⁻²		1	0	1			
$\mathrm{SD}<\!\!500~\mathrm{n}~\mathrm{mm}^{-2}$	0	1	0	0			
$SD > 1,000 \text{ n mm}^{-2}$	1	0	0	0			

- (5) SD The same method which was used for counting trichome density was used to determine SD.
- (6) Principal Coordinates analysis (PCoA) The PCoA is a method to explore and to visualize similarities or dissimilarities between data. The PCoA usually starts with a symmetric distance matrix and assigns for each item a location in a low-dimensional space [29, 30]. In this study, a taxa-characteristic matrix, including four fossil taxa plus the extant *Quercus delavayi*, and 14 characteristics were built and provided (Table 3). The PCoA was carried out using the MultiVariate Statistical Package (MVSP) software [31] and Euclidean distances after the data were standardized.

For descriptions of the leaf morphology and cuticles, we followed the terms of Dilcher [32], Luo and Zhou [33], and Ellis et al. [34].

2 Results

2.1 Systematics

Family: Fagaceae Hand.-Mazz., 1929Genus: *Quercus* L., 1753Subgenus: *Cyclobalanopsis* Oerst.

Str	ata	Thickness (m)	Stratigraphic column	Lithology	
Quaternary		1–20		Clay, silt, gravel	
		25.1		Light-gray thin-medium bedded mudstone bearing thin bedded silty mudstone	
		27.4		Light-gray medium-thick bedded silty mudstone	
		14.8		Dark-gray thin bedded silty mudstone	
		9.8	· N · · · N · · · N · · · N · · · N · · · N · · · N · · · N · · · N · · · N · · · N · · · N · · · N	Light-gary thin-medium bedded quartz sandstone	
		32.4	Image Image <th< td=""><td>Taupe thin-medium beeded mudstone, silty mudstone bearing lignite seams upper lignite seams 0.3 m, lower lignite seams 1.5–3m</td></th<>	Taupe thin-medium beeded mudstone, silty mudstone bearing lignite seams upper lignite seams 0.3 m, lower lignite seams 1.5–3m	
	_	9.7	• N • N • N • N • N • N • N • N • N • N • N • N • N	Prounosus medium-thick bedded feldspar quartz sand	
Pliocene	Ciying formation	47.1		Gray, grayish yellow medium bedded clay bearing silty mudstone	
		2.1	• N • N • N • N • N	Light-gray thin bedded fine-grained quartz sandstone	
		4.7		Light-gray spherullite mudstone bearling chalybeate	
			55.2		Light-gray thin-medium bedded ferritization silty mudstone
		23.4	Inc. Inc. <th< td=""><td>Light-gray thin-medium bedded ferritization silty mudstone</td></th<>	Light-gray thin-medium bedded ferritization silty mudstone	
		27.1		Gray-blue thin-medium bedded slity mudstone	
Oligocene Xiaotun formation				Prunosus thin-medium bedded mudstone	

No. Strata with fossils

Fig. 2 The sketch strata table of the Ciying Formation. The layer where the fossils were collected is marked with a *leaf symbol*. The figure is compiled from information from the Geological Bureau of Yunnan [21]

Species: Quercus tenuipilosa Q. Hu et Z.K. Zhou.

Etymology: The specific epithet *tenuipilosa* represents the rare trichomes of this species. In Latin, *tenuis* means "thin" or "weak", while *pilosa* means "hair" or "trichome" (Fig. 2).



Holotype: HST 856 A, B (counterparts) (Fig. 3a, b, e, and g) $\label{eq:stars}$

Paratype: HST 022 (Fig. 4c)

Other specimens: HST 151, HST 254, HST 565, HST 943A, and HST 943B (Fig. 4a, b, d-f)

Repository: All specimens examined are deposited in KUN.

2.2 Diagnosis

Leaves elliptic, symmetric; leaf apex acuminate and leaf base convex to broadly cuneate. Secondaries in the toothed portion craspedodromous, secondaries in the entire portion are camptodromous; tertiary veins opposite and alternate percurrent. Stomata on the abaxial epidermis, anomocytic. Trichome bases on the adaxial and abaxial epidermis, unicellular and multicellular.



Fig. 3 Comparison of the leaf morphology of *Quercus tenuipilosa* and the extant *Q. delavayi*. **a**, **b** *Quercus tenuipilosa*. Specimen no. HST856A, HST856B. *Scale bars* 1 cm. **c** *Q. delavayi*, DH037. *Scale bar* 1 cm. **d** Clear leaf of *Q. delavayi*. *Scale bar* 1 cm. **e** Close up of the upper portion of the fossil leaf. Specimen no. HST856A. *Scale bar* 1 cm. **f** Close up of the upper portion of the *Q. delavayi* leaf. *Scale bar* 0.5 cm. **g** Close up of the lower portion of the fossil leaf. Specimen no. HST856A. *Scale bar* 0.5 cm. **g** Close up of the lower portion of the fossil leaf. Specimen no. HST856A. *Scale bar* 0.5 cm. **b** Close up of the lower portion of the *Q. delavayi* leaf. *Scale bar* 0.5 cm.



Fig. 4 The paratype and other specimens of *Quercus tenuipilosa*. **a**-**f** HST943A, HST943B, HST022, HST151, HST254, and HST565. *Scale bars* 1 cm

2.3 Description

Gross morphology (Figs. 3, 4) Fossil blades are symmetric and elliptic in shape (Fig. 3a, b). The laminar size is 5.0 cm long, 2.0 cm wide (L:W ≈ 2.5). The blade apex is acuminate. The blade base shape is convex to broadly cuneate. The venation is pinnate with a straight and robust midvein. There are 13 pairs of secondary veins with regular spacing (Fig. 3a, b, e, and g). The major secondaries in the toothed portion are craspedodromous. The secondaries in the entire portion are camptodromous (Fig. 3e, g). The tertiary veins are both opposite and alternate, the outmost tertiaries are looped, running along the margin (Fig. 3e).

(2) Micromorphology (Fig. 5) The upper epidermis is composed of quadrangle to hexagonal cells. There are unicellular and multicellular trichome bases (Fig. 5a). The unicellular trichome bases are round, and the multicellular trichome bases are composed of 5–7 cells (Fig. 5a). Stomata are not found in the upper epidermis. The shape of cells in the lower epidermis



Fig. 5 Cuticular structures of *Quercus tenuipilosa* sp. nov. and *Q. delavayi* under the light microscope. **a** Adaxial epidermis of *Q. tenuipilosa*. Holotype. Slide No. HST856-20120814-upper-01. *Scale bar* 50 μm. **b** Adaxial epidermis of *Q. delavayi*. Slide No. DH030-3A-1. *Scale bar* 50 μm. **c** Abaxial epidermis of *Q. tenuipilosa*. Holotype. Slide No. HST856-20120814-lower-01. *Scale bar* 50 μm. **d** Abaxial epidermis of *Q. delavayi*. Slide No. DH030-4A-2. *Scale bar* 50 μm. **e** Anomocytic stomata of *Q. tenuipilosa*. Scale bar 10 μm. **f** Anomocytic stomata of *Q. delavayi*. Slide No. DH030-4A-2. *Scale bar* 10 μm. **g** Multicellular trichome base of *Q. tenuipilosa*. Slide No. HST856-20120814-lower-01. *Scale bar* 10 μm. **h** Multicellular trichome base of *Q. delavayi*. Slide No. DH030-4A-2. *Scale bar* 10 μm.

is irregular (Fig. 5c). The density of the multicellular bases is 250 n mm⁻² in the lower epidermis. The stomatal apparatuses are elliptical and randomly distributed in the areoles. A single ring of 5–8 subsidiary cells enclosing the two guard cells forms the ananomocytic type (Fig. 5c, g). The SD is ca. 672 n mm^{-2} .

3 Discussion

3.1 Systematic position of the fossil species

Jones [35] studied the leaf morphology and epidermis features of Fagaceae and related families and supplied a framework to identify fagaceous leaves. According to his results, our fossils differ from Juglandaceae by having symmetric bases and differ from Betulaceae by the absence of doubly serrate margins. The synthetic morphologic characteristics of our fossils, such as the symmetric leaf base, possessing unicellular trichome bases, serrate but not doubly serrate, clearly place our fossils into Fagaceae. In the Fagaceae, only Lithocarpus and the Quercus subgenus Cyclobalanopsis possess the same secondary veins as our fossils [33]. However, *Lithocarpus* leaves are usually entire or only partly toothed [35], and most *Lithocarpus* possess a cuneate leaf base rather than the convex base [14]. In China, only eight Lithocarpus species are serrated [14], but they differ from our fossils by having larger leaves, with many more secondaries. Thus, our fossils are assigned to the Quercus subgenus Cyclobalanopsis. We compared our fossils with all extant Cyclobalanopsis species in China using the characters listed in Deng [36] and Xing et al. [19]. We found that our fossils are similar to Q. glauca, Q. schottkyana, and Q. delavayi. However, the fossil species has both unicellular and multicellular trichome bases, while Q. glauca and Q. schottkyana only have unicellular trichome bases on the abaxial epidermis [19]. Thus, the fossil species is most similar to the extant Q. delavayi.

Until now, several fossil species have shown a close affinity with the extant Q. delavayi including Q. aff. delavayi from the late Miocene Tiantai flora, Zhejiang Province [37] and Q. praedelavayi from the late Miocene Xianfeng flora, Yunnan Province [19]. Among these fossil species, Q. aff. delavayi possesses much narrower leaves than our fossils with L:W > 5, and Q. praedelavayi is very similar to our fossils in the leaf architecture but differs from our fossils by possessing larger leaves and lacking trichome bases on the adaxial epidermis (Table 2). Q. praedalavayi, Q. aff. delavayi, and our fossils share similar epidermis characteristics such as cell shape and trichome base type. However, our fossils and Q. aff. delavayi possess both unicellular and multicellular trichome bases on the adaxial epidermis, while Q. praedelavayi does not possess trichome bases on the adaxial epidermis [19]. Q. aff. delavayi and Q. praedelavayi also differ from our fossils in the trichome and stomatal densities (Table 2). To explore the similarities between our fossils and the other two fossil species and the extant Q. delavayi, we carried out the PCoA using 14 characteristics (Table 3; Fig. 3). The results indicated that Q. aff. delavayi is distinct from the other three species. Our fossils are most similar to the extant Q. delavayi, followed by Q. praedelavayi. Considering the differences in trichome and stomatal densities, we describe it as a new fossil species, Q. tenuipilosa. As Q. tenuipilosa and Q. praedelavayi both show close affinity with Q. delavayi, they could be considered as part of the Q. delavayi complex.

3.2 Evolution of the trichome density of the *Quercus delavayi* complex

Epidermal trichomes, sunken stomata, and stomatal furrows are commonly considered as adaptations to aridity, increasing the boundary layer resistance and consequently



PCO case scores (Gower General Similarity Coefficient)

Fig. 6 The morphologic principal coordinates analysis (PCoA) based on 14 characteristics of Quercus tenuipilosa and related species



Fig. 7 The trichome density and SD of the Quercus delavayi complex

limiting transpiration [1, 2, 38, 39]. However, many plants in humid regions also have trichomes [40]. Previous studies have indicated that trichomes may have other functions, such as anti-herbivory defenses, reducing incident light, and decreasing photosynthetic rates [41, 42].

Hardin studied the seasonal trichome variations of *Quercus* from Eastern North America and found the trichome density of *Quercus* was affected by environmental change [43]. Levin demonstrated that the morphology and density of trichomes show linear relationships with environmental parameters, but the responses are species-specific [44]. In this study, we showed that the trichome densities of the *Q. delavayi* complex decreased from the Late Miocene to the present epoch (Table 1; Fig. 4). Palaeoclimatic reconstructions suggest that the climate during the Late Miocene in Yunnan was warmer and more humid and became cooler and drier since the Pliocene [45–51]. The relationship between the trichome density of the *Q. delavayi* complex and the palaeoclimate contradict the



Fig. 8 The regression of the trichome densities of the extant *Quercus* delavayi in relationship to the mean annual precipitation of 27 localities

hypothesis that leaf trichomes are an adaptation to the cooler and drier environment [8, 52]. To further test this hypothesis, we studied the trichome densities of extant Q. *delavayi* from 27 localities and made a regression against the mean annual precipitation. Analysis of the regression results indicated that there is no significant relationship between the trichome density of Q. *delavayi* and local precipitation ($R^2 = 0.0579$, P = 0.1196). Haworth et al. [53] also showed that xeromorphic traits are not restricted to plants subjected to water stress but also serve multiple functions such as water-repellence, defense, and protection



Fig. 9 The Cenozoic CO₂ concentration and the stomatal densities of the Quercus delavayi complex. P Pliocene

from excess light. Further ecologic and growth experiments are required to fully understand the function of epidermis trichomes.

3.3 Evolution of the SD of the *Quercus delavayi* complex

As a gateway for gas exchange and transpiration of plants, SD, size, and conductance are sensitive to changes in the CO_2 concentration [7, 53–55]. The relationship between SD, stomatal index (SI), and CO₂ concentration change has been investigated intensively, showing either positive or negative responses (e.g., [54, 56]). Royer analyzed the responses of the SD and SI of 176 C3 plants to CO2 concentration. The results showed that 40 % and 36 % of the species (for SD and SI, respectively) in experimental studies and 88 % and 94 % of the species in fossil studies showed inverse relationships, while <12 % of the species show positive relationships with CO_2 concentration [56]. Haworth measured the SI of six Cupressaceae species and found that the SI of three species (Tetraclinis articulata, Callitris columnaris, and C. rhomboidea) showed significant inverse relationships with rising CO₂, while the other three species (Athrotaxis cupressoides, C. preissii, and C. oblonga) showed no response in SI [53] (Figs. 6, 7, 8).

Furthermore, to test whether CO_2 concentration is the main factor influencing SD and SI, Kouwenberg et al. [8] and Royer [56] discussed the impact of other climatic parameters (temperature, precipitation, and UV-B radiation) on the stomatal frequency, respectively. Their results

showed that CO_2 is the most important factor influencing the SD and SI.

In this study, we determined the SD of the Quercus delavayi complex from the Late Miocene, the Late Pliocene, and the present epoch (Table 2; Fig. 4). The results indicated that the SD of the Q. delavayi complex was the highest during the late Miocene, declined in the late Pliocene, and then increased again at the present epoch. The Cenozoic CO2 concentrations from these three epochs show the opposite trend. The CO_2 concentration was the lowest in the Late Miocene, increased dramatically during the Pliocene, and dropped again until the present epoch (pre-industrial) [57]. These results indicated that the SD of the Q. delavayi complex showed an inverse relationship with atmospheric CO₂ concentrations, suggesting that the SD of the Q. delavayi complex may be a good proxy for reconstruction of paleo-CO₂ concentrations (Fig. 9).

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