Contents lists available at ScienceDirect







journal homepage: www.elsevier.com/locate/ecolind

# Influence of climate factors on the global dynamic distribution of *Tsuga* (Pinaceae)

Shumei Xiao<sup>a,b</sup>, Shufeng Li<sup>a,\*</sup>, Jian Huang<sup>a</sup>, Xiaojun Wang<sup>b,c</sup>, Mengxiao Wu<sup>d</sup>, Rizwan Karim<sup>a</sup>, Weiyudong Deng<sup>e</sup>, Tao Su<sup>a,\*</sup>

<sup>a</sup> CAS Key Laboratory of Tropical Forest Ecology, Xishuangbanna Tropical Botanical Garden, Chinese Academy of Sciences, Mengla 666303, Yunnan, China

<sup>b</sup> University of Chinese Academy of Sciences, Beijing 100049, China

<sup>c</sup> CAS Key Laboratory of Cenozoic Geology and Environment, Institute of Geology and Geophysics, Chinese Academy of Sciences, Beijing 100029, China

<sup>d</sup> Senckenberg Natural History Collections Dresden, Königsbrücker Landstraße 159, 01109 Dresden, Germany

<sup>e</sup> Rheinische Friedrich-Wilhelms-Universität Bonn, 53115 Bonn, Germany

# ARTICLE INFO

Keywords: Tsuga Climate change Species conservation MaxEnt model Pollen

# ABSTRACT

Throughout the Quaternary period, climate change has significantly influenced plant distribution, particularly affecting species within the genus Tsuga (Endl.) Carrière. This climatic impact ultimately led to the extinction of all Tsuga species in Europe. Today, there are ten recognized species of Tsuga worldwide, one of listed as a vulnerable species and four as near-threatened species. The genus Tsuga exhibits a disjunctive distribution in East Asia (EA), eastern North America (ENA), and western North America (WNA). It is crucial to comprehend the mechanisms underlying these distributional changes and to identify key climate variables to develop effective conservation strategies for Tsuga under future climate scenarios. In this study, we applied the maximum entropy (MaxEnt) model by combining distribution data for Tsuga with abundant pollen fossil data. Our objective was to investigate the climate factors that shape the distribution of Tsuga, identify climate thresholds, and elucidate distribution dynamics in the context of significant climate changes over the past 1070 thousand years (ka). Our findings highlight the pivotal role of precipitation as the key climate factor affecting the distribution of Tsuga. Specifically, in EA, summer precipitation was the key driver, while in North America (NA), winter precipitation exerted greater importance. Moreover, we observed similarities in climatic requirements between *Tsuga* species in Europe and EA, and declines in summer precipitation and winter temperature were major factors contributing to the extinction of Tsuga species in Europe. Quaternary glacial and interglacial fluctuations exerted substantial impacts on Tsuga distribution dynamics. The disappearance of Tsuga species in the Korean Peninsula may have occurred during the LGM (Last Glacial Maximum). The potential suitable area for Tsuga species in EA expanded during the cold periods, while in NA, it contracted. In the future, climate change may result Tsuga distribution area contraction in both the EA and NA. Our study has identified distinct response patterns of Tsuga in various geographic regions to Quaternary climate change and offers corresponding suggestions for Tsuga conservation. In the future, it will be imperative to prioritize the conservation of natural Tsuga distributions in EA and NA, with a focus on the impacts of precipitation fluctuation on the dynamic distribution of this genus.

#### 1. Introduction

The Quaternary climate experienced repeated temperature fluctuations (Hewitt, 2003), mainly manifested in the periodic growth and retreat of continental glaciers in the Northern Hemisphere (Hansel and McKay III, 2010; Willeit et al., 2019). The severe climate changes had profound impacts on the migration, diffusion, differentiation, and even extinction of species (Van Andel and Tzedakis, 1996; Sandel et al., 2011), leading to dramatic biodiversity changes (Davis, 1983; Hewitt, 2000) and the development of a fragmented distribution pattern among certain species in the Northern Hemisphere (Hong et al., 2018). Compared to other geological periods, the Quaternary was more closely related to the modern era and therefore had a significant impact on current Northern Hemisphere biodiversity (Zhou et al., 2017). Understanding climate change and species response mechanisms in the Quaternary is of great significance for gaining a deeper understanding of

\* Corresponding authors. E-mail addresses: lisf@xtbg.org.cn (S. Li), sutao@xtbg.org.cn (T. Su).

https://doi.org/10.1016/j.ecolind.2023.111533

Received 23 August 2023; Received in revised form 12 December 2023; Accepted 18 December 2023 Available online 6 January 2024

1470-160X/© 2023 The Author(s). Published by Elsevier Ltd. This is an open access article under the CC BY-NC-ND license (http://creativecommons.org/licenses/by-nc-nd/4.0/).



global climate change and ecosystem relationships (Fordham et al., 2020), as well as for predicting the impact of future climate change on species distributions and the preservation of threatened species (Fordham et al., 2014, 2016; Bhattacharyya et al., 2023).

Tsuga (Endl.) Carrière, in the family Pinaceae, plays a significant role in subalpine and lowland moist coniferous forests as well as broad--leaved forests (Lepage, 2003). Tsuga have been greatly affected by Quaternary climate change (Magri and Palombo, 2013), leading to its disappearance in Europe (Davis, 1983; Magri et al., 2017). Currently, Tsuga have a fragmented distribution pattern in East Asia (EA) and North America (NA) (Li, 1995; Lepage, 2003; Qiao et al., 2007; Holman et al., 2017). The widely accepted classification contains nine extant species, including T. chinensis (Franch.) Pritz., T. forrestii Downie and T. dumosa (D. Don) Eichler distributed in China and other countries around the Himalayas. T. diversifolia (Maxim.) Mast. and T. sieboldii Carrière are endemic in Japan, T. canadensis (L.) Carrière, T. caroliniana Engelm., T. heterophylla (Raf.) Sarg. and T. mertensiana (Bong.) Carrière are distributed in NA (Farjon, 2017). In addition, a new species, T. ulleungensis G. P. Holman et al., has recently been reported and distributed on Ulleungdo Island, Korea (Holman et al., 2017; Feng et al., 2021). Among them, T. forrestii is listed as vulnerable species (Yang and Luscombe, 2013, Yang et al., 2017), and some others are listed as nearthreatened species including T. sieboldii (Katsuki and Luscombe, 2013), T. canadensis (Farjon, 2013a), T. caroliniana (Farjon, 2013b), T. chinensis var. robusta and T. mertensiana (Yang et al., 2017). Therefore, assessing the distribution status of Tsuga has important significance for the conservation of *Tsuga*, especially considering in the future climate scenarios.

This genus thrives in shaded and humid environments and is sensitive to prolonged droughts, it is usually considered an excellent indicator of humid environment in the geological time (Lepage, 2003; Yang et al., 2009; Lacourse, 2012; Xing et al., 2013; Ding et al., 2021). The macrofossil occurrence of *Tsuga* is frequently used to indicate warm and humid climate in the geological time, such as the Miocene climate in Weichang, Hebei Province, North China (Li et al., 2023); the late Paleogene climate in Markam, southeastern margin of the Qinghai–Tibetan Plateau (Wu et al., 2020); the Miocene climate in Xundian County, Yunnan Province, southwestern China (Wang et al., 2015). The above research indicates that *Tsuga* is a good indicator of climate change and has received wide attention. However, the most important climate variables driving the distribution of *Tsuga* are not yet clearly identified. Therefore, conducting research on the distribution dynamic of *Tsuga* has important environmental inductive significance.

Tsuga is commonly found in fossil records (Lepage, 2003; Magri et al., 2017), particularly in the form of pollen grains. Its abundant Quaternary pollen fossils make it an exceptional conifer plants for exploring the impact of Quaternary climate change on its species distribution. Magri et al. (2017) observed the disappearing pattern of Tsuga in southern Europe during the Quaternary, and Davis et al. (1986) suggested an expansion route of Tsuga in the USA around 7 thousand years (ka). In addition, Bertini (2000) found that extinction of Tsuga in Europe was associated with global drought events during the middle Pleistocene, and Prentice et al. (1991) indicated that a cool and humid climate promoted the expansion of T. mertensiana during 12-9 ka. Although previous palynology studies have revealed the distribution pattern and key climate variables of Tsuga during the Quaternary, they just contained part of the region and species. In this regard, species distribution modeling in palaeoecology is a valuable approach for predicting past species distributions in large spatial scale and continuous time series (Svenning et al., 2011; Varela et al., 2011) and examining the relationship between climate change and suitable species habitats (Elith et al., 2011). However, previous research using models to simulate the distribution of Tsuga are constrained in some areas such as Japan (Tsuyama et al., 2014), southwestern China (Dakhil et al., 2019), Maine, U.S.A (Dunckel et al., 2017), and limited species. There is still lack of integrative research on its Quaternary distribution dynamic of all species in Tsuga, the specific climate factors that contributed to the extinction of Tsuga in Europe is not yet fully understand.

The MaxEnt model, also known as the maximum entropy model, is based on the principle of maximum entropy theory (Phillips and Dudík, 2008). It uses existing species distribution and environmental data (Phillips et al., 2006; Elith et al., 2011) to estimate a species distribution by identifying the distribution that exhibits maximum entropy constrained by the environmental conditions observed at recorded locations (Phillips et al., 2006, 2017). Compared with other models, MaxEnt offers notable advantages such as high simulation accuracy and satisfactory modeling results, even with relatively small sample sizes (Guo et al., 2020). Consequently, it has been widely applied across various ecological research fields (Phillips et al., 2017; Liu et al., 2019; Xiao et al., 2022).

According to Magri et al. (2017), the distribution pattern of *Tsuga* in Southern Europe changed dramatically after 1070 ka. Therefore, this study used MaxEnt to simulate the global distribution change of *Tsuga* from 1070 ka into the future and analyze critical bioclimate variables. Additionally, pollen fossil data were collected from previous publications and compared with the modeling outputs to facilitate mutual validation and comprehensive analyses. This study aimed to address three key scientific questions: 1) What was the impact of Quaternary climate change on the distribution of *Tsuga* in Europe; and 3) What are potential distribution regions of *Tsuga* under future climate change. We hypothesize that precipitation and temperature in different seasons may affect the distribution and extinction of *Tsuga* in various regions.

#### 2. Materials and methods

#### 2.1. Study regions

This study used data on the global natural distribution of *Tsuga*, which exhibits a fragmented pattern in EA, eastern North America (ENA), and western North America (WNA, Fig. 1). In EA, *Tsuga* are mainly distributed in the Asian monsoonal regions of Japan, China, and other countries around the Himalaya (Yang et al., 2009). In ENA, *Tsuga* are distributed in the eastern United States and Canada, characterized by a humid and cold climate with sufficient water throughout the year (Hough, 1960). In WNA, *Tsuga* thrive in cool to cold marine climates characterized by mild to cold winters, short and warm to cool growing seasons, and moderate to heavy precipitation (Berntsen, 1958; Means, 1990). Due to the distinct differences in climatic and environmental conditions in these three regions, the impacts of Quaternary climate change on *Tsuga* were significantly varied different (Eiserhardt et al., 2015). Therefore, this study conducted separate simulations individually for these three regions.

# 2.2. Species data collection and processing

The extant distribution data of *Tsuga* came from online sources, namely the Chinese Virtual Herbarium (https://www.cvh.ac.cn/) and the Global Biodiversity Information Facility (https://www.gbif.org/; GBIF.org, 2022). A total of 46,185 modern distribution records were collected. To ensure data quality, a rigorous cleaning process was conducted, which excluded missing location information, duplicate entries, and cultivated occurrences. After data cleaning, 36,842 records remained. To avoid spatial bias, the remaining records were randomly gridded using the 'grid Sample' function from the 'dismo' package in the software R. Only one occurrence per  $20' \times 20'$  area was retained, resulting in a final dataset of 2,425 standard records used for distribution modeling (Fig. 1).

In addition to the modern distribution data, a collection of pollen fossil records was compiled. A total of 268 pollen fossil records were collected from various sources (Table S1), including studies by the relevant references (McLachlan and Brubaker, 1995; Jacques et al., 2000; Toney et al., 2003; Hayashi, 2010; Xu et al., 2010; Tarasov et al.,



Fig. 1. Modern distribution of Tsuga at the global scale (a), NA (b) and EA (c).

2011; Corrado and Magri, 2011; Sun and Feng, 2015; Donders et al., 2021; Liu and Yuan, 2022 etc.), and two pollen databases, Neotoma Explorer (https://apps.neotomadb.org/explorer/) and Paleobiology Database (https://palaeobiodb.org/).

# 2.3. Bioclimate variables and selection

Climate data were primarily obtained from the World Climate Database (https://www.worldclim.org) with a spatial resolution of 2.5 arc minutes (~5 km) and encompassing 19 bioclimate variables (Table S2). The climate data cover multiple periods, including the Last Glacial Maximum (LGM, ca. 22 ka BP, CCSM4), the mid-Holocene (MH, ca. 6 ka BP, CCSM4), the present period (average for 1970-2000), and the future (2021-2040 and 2041-2060). Future climate data were derived from two shared socioeconomic pathways (SSP245 and SSP585) of five models (ACCESS-CM2, CMCC-ESM2, EC-Earth3-Veg, MRI-ESM2-0, UKESM1-0-LL) from the Coupled Model Intercomparison Project Phase 6 (CMIP6). The Quaternary climate data preceding the LGM were derived from Oscillayers (https://doi.org/10.5061/dryad.27f8s90, Gamisch, 2019). These data were chosen based on the global mean surface temperature (Gamisch, 2019, refer to Fig. S1) and the consideration of the dynamic changes in pollen fossil data. The selected Quaternary climate data included periods of 1,070 ka, 880 ka, 780 ka, 440 ka, LGM and MH. Of these, 1,070 ka, 780 ka, and MH represented warm periods, while 880 ka, 440 ka, and LGM represented cold periods.

To reduce the impact of overfitting due to climate variable multicollinearity, we used the "laverStats" function from the package "spatialEco" in R to analyze Pearson's correlation coefficient of bioclimate variables. Bioclimate variables with high contributions and low correlations (<0.8) were selected (Table S3, S4). Consequently, in EA, we selected bio6 (min temperature of coldest month/°C), bio7 (temperature annual range/°C), bio9 (mean temperature of driest quarter/°C), bio10 (mean temperature of warmest quarter/°C), bio11 (mean temperature of coldest quarter/°C), bio15 [precipitation seasonality (coefficient of variation)], bio18 (precipitation of warmest quarter/mm) and bio19 (precipitation of coldest quarter/mm). In ENA, we selected bio7, bio9, bio10, bio11, bio12 (annual precipitation/mm), bio15, bio17 (precipitation of driest quarter/mm) and bio18 as relevant variables. In WNA, we selected bio3 [isothermally (bio2/bio7) (×100)], bio4 (temperature seasonality), bio8 (mean temperature of wettest quarter/°C), bio10, bio11, bio15, bio18 and bio19.

# 2.4. Species distribution model building

We used the species distribution model MaxEnt (version 3.4.3, https ://biodiversityinformatics.amnh.org/open\_source/MaxEnt/) to simulate the potential distribution of *Tsuga* in the EA, ENA, and WNA regions under different climate scenarios. MaxEnt is efficient when handling presence-only data and species-environment relationships (Elith et al., 2011; Phillips et al., 2006). Setting up the model involved randomly selecting 75 % of the distribution data to be used as training data, while the remaining 25 % was treated as testing data. This process was repeated 10 times using a cross–validation procedure. The area under the receiver operating characteristic curve (AUC value) was used to evaluate model performance. AUC values range from 0.5 to 1.0, where AUC > 0.9 indicates better model performance (Phillips and Dudík, 2008). We conducted a jackknife test to assess the relative importance of all variables.

The MaxEnt predictions were converted into raster data. We then classified the raster data using a reclassification method proposed by Cao et al. (2022), in which the potential adaptability of *Tsuga* was divided into four suitability ranges: unsuitable (0.0-0.1), low suitability (0.1-0.3), medium suitability (0.3-0.5), and high suitability (0.5-1.0). These classifications provided insights into the changes to potentially suitable regions for *Tsuga* in response to different climate scenarios. Then we use subtraction tools to calculate changes in climate scenarios and potential distribution regions.

# 3. Results

#### 3.1. Model performance, variable contributions, and response curves

The model simulation results showed that AUC values for EA (0.977), ENA (0.948), and WNA (0.973) were all greater than 0.9 (Fig. S2). These high AUC values indicated that the simulations obtained in this study were highly reliable, demonstrating the robustness of our approach. The contribution rates of the eight main bioclimatic variables in the three regions are shown in Fig. 2. Overall, precipitation was identified as the



Fig. 2. The contribution rates of several important climate variables to the modern distribution of Tsuga in the EA (a), ENA (b), and WNA (c) regions.

crucial climate variable influencing the distribution of Tsuga in all regions. However, the influence of specific precipitation variables varied across regions. In EA, the precipitation of the warmest quarter (bio18) in summer had a significant effect on the distribution of Tsuga. In ENA, the precipitation of the driest quarter (bio17) in winter played an important role in species survival. Similarly, in WNA, the precipitation of the coldest quarter (bio19) during winter had a notable impact on the distribution of Tsuga. Because the driest quarter in ENA and the coldest quarter in WNA both occurred during winter, it was inferred that winter precipitation played a critical role in the survival and distribution of Tsuga in NA. In conclusion, summer precipitation had a crucial influence on the distribution of Tsuga in EA, while winter precipitation played a significant role in NA. Additionally, annual precipitation (bio12) impacted the distribution of Tsuga in the ENA, and precipitation seasonality (bio15) and annual temperature range (bio7) affected the two regions. Both winter and summer temperatures had important effects on the survival and distribution of Tsuga in all regions.

The response curves provide valuable insights into species ecological

niches and suitable bioclimatic variable ranges for its habitat (Liu et al., 2021). Fig. 3 showed the response curves of Tsuga in EA, ENA, and WNA for important climates. The high suitability area was defined within a range of 0.5–0.1. The area with a response curve greater than 0.5 was considered to fall within the threshold range of climate variables suitable for Tsuga survival. In EA, the response curves showed that Tsuga had the strongest response to precipitation during the warmest season, particularly when it exceeded 500 mm. Additionally, the mean temperature of the coldest quarter ranging from 0 to 10 °C had a significant influence on Tsuga habitat suitability. Similarly, in ENA, annual precipitation (900-1400 mm) and annual temperature range (34-41 °C) produced strong impact on Tsuga distribution. These climatic variables played a crucial role in determining the suitability of habitats for Tsuga in the ENA region. In WNA, the important variables included precipitation of the coldest quarter (440-1750 mm), mean temperature of the wettest quarter (-6-6 °C), and mean temperature of the warmest quarter (12-17 °C), which had a strong influence on the distribution and viability of Tsuga.



Fig. 3. Response curves of Tsuga to important climate variables in EA (a-c), ENA (d-f), and WNA (g-i).

# 3.2. Comparison of the predicted and actual Tsuga distribution regions

The simulated model results from this study closely matched the natural distribution sites of Tsuga. In EA, the potential suitable area was basically consistent with the natural distribution regions, including the surrounding regions of the Sichuan Basin, Qinling Mountains, Hengduan Mountains, Himalaya, Hills in Southeast China, Central Mountain Range on the island of Taiwan, and Honshu, Shikoku, and Kyushu in Japan (Fig. 4a-b). The estimated potential high suitable areas in EA covered 2.85 million km<sup>2</sup> (Fig. S3a). Interestingly, there were small regions of potential high suitable outside the natural distribution regions, such as the Taebaek Mountains and southeastern Korean Peninsula, and parts of Jeju Island and the European Alps (Fig. 4b). In ENA, the species was primarily distributed in the Appalachian Mountains, the Great Lakes region (Fig. 4c-d). The estimated potential high suitable areas in ENA spanned 3.93 million km<sup>2</sup> (Fig. S4a). In WNA, the main distribution regions for Tsuga included the Coast Mountains, the Rocky Mountains, and other coastal regions (40-55 °N, Fig. 4e-f), with a potential high suitable area of 2.63 million km<sup>2</sup> (Fig. S5a). Notably, some areas had potential medium suitable for the fragmented distribution of Tsuga in Europe (Fig. 4f). In conclusion, the potential suitable regions for modern Tsuga based on the natural distribution areas in EA and ENA were generally consistent with the modern natural distribution. Additionally, the study suggested there may be some potential suitable areas for Tsuga in Europe beyond its current distribution range (Fig. 4).

# 3.3. The potential suitable regions of Tsuga in the past

The study simulated the past potential distribution regions of *Tsuga* in EA, ENA, and WNA and compared them with collected pollen fossil data (Fig. 5–Fig. 7). The pollen fossil data were mainly concentrated in the LGM and MH periods, and they were found within or surrounded the regions of potential high suitable (Fig. 5e–f, Fig. 6e–f, Fig. 7e–f). Compared with previous pollen research results by Magri et al. (2017), this study found that the distribution of pollen sites in Europe was generally consistent with the simulation results of *Tsuga* in EA, with *Tsuga* disappearing in Europe after 780 ka (Fig. S6–S8). Consequently, pollen sites in Europe were added to the simulated map of the potential distribution of *Tsuga* in EA (Fig. 5A). Overall, the distribution of pollen sites corresponded well with the potential suitable regions, suggesting that the simulation results effectively represented the dynamic distribution change of *Tsuga* during geological periods.

In EA, the area of potential suitable for *Tsuga* during warm periods was significantly smaller than that during cold periods (Fig. 5a-f). The LGM exhibited the largest distribution area of potential suitable regions for Tsuga (Fig. 5e). The periods from 1,070 to 880 ka and from 780 to 440 ka showed the most significant increase in potential suitable regions for Tsuga, with the area of potential high and medium suitable expanding by 86.56 %, 60.14 %, 28.17 %, and 37.88 %, respectively (Fig. S3a-b). The expansion areas were primarily located in the Himalaya, southern Hengduan Mountains, Hills in Southeast China, Central Mountain Range on the island of Taiwan, and southern Japan, while the contraction areas were mainly concentrated in middle and northern Japan (Fig. 5a-d). The potential distribution areas for Tsuga disappeared in the Korean Peninsula during the LGM (Fig. 5e). The spatial distribution of potential suitable regions shifted notably to lower latitudes and elevations (Fig. 5e). The most obvious reduction in the potential suitable area of Tsuga occurred during 880-780 ka and from the LGM to MH. The potential high and medium suitable area decreased by 30.43 % and 54.23 % and by 18.36 % and 14.65 %, respectively (Fig. S3a-b). The contracting regions were primarily located in the Himalayas, southern Hengduan Mountains, Hills in Southeast China, Central Mountain Range on the island of Taiwan, and southern Japan. The spatial distribution of potential suitable regions moved significantly to higher latitude and elevation areas (Fig. 5b-c, e-f). The expanding regions were mainly distributed in the central area of Honshu Island, the southern Ou Mountains in Japan, and the Korean Peninsula (Fig. 5b–c, e–f). In summary, throughout geological periods, the potential habitat of *Tsuga* experienced both expansions and contractions in response to climate change. The regions where the potential habitat of *Tsuga* contracted were largely congruent with the regions where it expanded, indicating that *Tsuga* populations in these regions were highly sensitive to climate change.

In ENA, the potential suitable area for *Tsuga* was considerably larger during relatively warm periods than during relatively cold periods (Fig. 6). The smallest potential suitable area for Tsuga were observed during the LGM. The most pronounced contractions in the potential suitable area for Tsuga occurred during 1,070-880 ka and 780-440 ka, with areas of potential high and medium suitable decreasing by 40.38 %, and 48.27 %, 51.8 % and 46.52 %, respectively (Fig. S4a-b). Reductions in regions were primarily distributed in the Great Lakes, central and northeastern Appalachian Mountains, and northern Newfoundland (Fig. 6a-d). On the other hand, the most significant expansions in potential suitable area for Tsuga were observed during 880-780 ka, LGM-MH, and MH-Present, with areas of potential high suitable expanding by 53.82 %, 81.31 %, and 39.46 %, respectively (Fig. 4a). The potential medium suitable areas expanded by 38.65 % (880-780 ka) and 177.12 % (LGM-MH) (Fig. S4b). The expanded regions were mainly distributed in the Great Lakes, central Appalachian Mountains, with potential suitable noticeably shifting to higher latitudes and elevations (Fig. 6b-c, e-f).

In WNA, the potential suitable area for Tsuga was larger during warm periods than during cold periods (Fig. 7). The significant contractions in areas of potential high suitable mainly occurred during 1,070-880 ka and 780-440 ka, with reductions of 36.72 % and 27.23 %, respectively (Fig. S5a). The potential high suitable area retreated southwards along the Coast Mountains and Rocky Mountains, extending even to regions south of 45.5 °N (Fig. 7a-d). Some of the regions classified as high suitable regions transformed into medium and low suitable regions. On the other hand, the significant expansion periods of potential high suitable area for Tsuga occurred mainly in 880-780 ka, LGM-MH, and MH-Present, with expansions of 48.17 %, 20.58 %, and 60.55 %, respectively (Fig. S5a). The potential high suitable area expanded northwards and eastwards along the Coast Mountains and Rocky Mountains, extending further to 58 °N (Fig. 7b-c, e-f). Parts of the medium and low suitable regions were transformed into high suitable regions that expanded inland and into higher latitudes.

# 3.4. Analysis of changes in Europe based on important climate variables

Palynological studies have found that there was a substantial contraction of Tsuga populations in Europe during the late early Pleistocene, followed by a gradual disappearance in the middle Pleistocene (Biltekin et al., 2015; Magri et al., 2017). The decline and eventual disappearance of Tsuga populations in Europe during the middle Pleistocene may have been associated with intensified drought and decreased temperatures during glacial periods (Bertini, 2000). This study considered the important climate variables affecting the survival and distribution of Tsuga in EA (summer precipitation and winter temperature) and NA (precipitation of coldest quarter and annual precipitation), and further examined the changes to these climate variables in Europe (Fig. 8). The changes in the most influential climatic factors in EA were consistent with distribution shifts of Tsuga pollen in Europe, while NA was different (Fig. S6-S8). Specifically, there was a transition from a relatively warm period to a cold period, the percentage of Tsuga pollen in Europe decreased, and the corresponding climate variables (summer precipitation and winter temperature) significantly decreased (Fig. 8a1, a3-a4; c1, c3-c4). However, the critical climate variables for the distribution of Tsuga in NA did not correspond to the distribution changes in Europe. These findings suggest that Tsuga in Europe may have had more similar climate requirements to those in EA, and reductions to summer precipitation and winter temperatures may have caused its



Fig. 4. Comparing modern natural and potential distribution regions of Tsuga in EA (a-b), ENA (c-d) and WNA (e-f).



Fig. 5. Predict past potential distribution regions of Tsuga in EA.











Fig. 8. The variations in four important climate variables in Europe. Note: these four climate variables were found to be the most important climate variables affecting the survival and distribution of *Tsuga* in EA (bio18 and bio11) and NA (bio19 and bio12).

disappearance.

### 3.5. Potential suitable regions for future Tsuga

In EA, the areas of future potential suitability for Tsuga were projected to contract even further, with a shift towards higher latitude and elevation regions. In particular, under the higher carbon emission scenarios (SSP585), the potential suitable area decreased significantly (Fig. S3a-b) and moved towards higher elevations in regions like those in the surrounding Himalaya, southwestern Qinling Mountains, Hengduan Mountains, and the Kyushu and Shikoku Islands in Japan (Fig. S9b, d). In ENA, the potential suitable area for Tsuga was projected to decrease in the future, especially under SSP585 (2041-2060), with a noticeable reduction in the distribution area (Fig. S4a). The decreased areas were projected to be mainly distributed in the southwest Appalachian Mountains, while expansions were predicted to occur in the areas north of the Great Lakes in NA and Labrador Plateau. (Fig. S10d). In general, under the high carbon emissions scenario, the potential suitable area for Tsuga noticeably expanded to higher latitudes and altitudes. In WNA, the potential suitable area for Tsuga in the future was expected to generally decrease compared to present conditions (Fig. S5, Fig. S11). Under future climate scenarios, the potential suitable area significantly contracted, especially in SSP585(2021–2040) and SSP585 (2041–2060) climate scenarios, with high suitable area decreasing by 26.37 % and 35.03 %, respectively (Fig. S5a). These areas were predominantly located in the Coast Mountains (45–48 °N) and the southern Rocky Mountains (Fig. S11b, d).

# 4. Discussion

#### 4.1. Key bioclimatic variables influencing Tsuga distribution

This study highlighted the significant impact of climate change on the distribution of *Tsuga* (Fig. 3, Fig. S12). Precipitation was the most crucial climate variable affecting *Tsuga* distribution, with different regions showing sensitivity to either summer or winter precipitation. Previous research has shown that *Tsuga* prefer humid climates and exhibit limited drought tolerance (Lepage, 2003; Yang et al., 2009). Dry climatic conditions in the geological past have been linked with the disappearance of *Tsuga* various regions of EA and NA (Lepage, 2003), such as in Markam, the southeast margin of the Qinghai-Tibetan Plateau (Wu et al., 2020); Xianfeng Basin, central Yunnan Wang et al. (2015); Hokkaido, Japan (Tsukada, 1985); and interior Alaska, WNA (Ager et al., 1994). Specifically, in EA, Tsuga needs high levels of precipitation for survival, especially in summer (Yang et al., 2009; Tsuyama et al., 2014; Ni et al., 2019). In NA, the Tsuga distribution area needs high annual and winter precipitation. Research has found that 90 % of the Tsuga population in NA is distributed in areas where the annual precipitation exceeds 800 mm (Thompson et al., 1999). Increased precipitation in the geological past has expanded the distribution range of Tsuga (Calcote, 2003; Williams et al., 2004; Munoz et al., 2010). Adequate water availability in winter is crucial for the seed germination and seedling growth of Tsuga (Godman and Lancaster, 1990; Means, 1990), and drought conditions can lead to seed mortality (Berntsen, 1958). Our research showed that Tsuga had distinctive seasonal precipitation demands in different regions.

Furthermore, the survival of Tsuga was strongly influenced by winter and summer temperatures (Fig. 3, Fig. S12). The annual temperature range and precipitation seasonality impacted the survival and distribution of Tsuga in EA and ENA (Fig. 3, Fig. S12). Previous studies have indicated that vernalization, a process requiring exposure to cold temperatures, is necessary for seed germination of *Tsuga* spp. (Vasiliauskas, 1996). In EA, Tsuga are naturally distributed in the monsoon region (Yang et al., 2009). It is sensitive to winter and annual temperature ranges (An et al., 2011). In NA, Tsuga growth necessitates approximately 10 weeks of vernalization at temperatures slightly below or around the freezing point to achieve the highest germination rate (Godman and Lancaster, 1990). In its natural distribution range, the average temperature is -8 °C in January, while it reaches 21 °C in July (Oswald and Foster, 2012). A decrease in average January temperatures leads to a contraction in the distribution range of Tsuga (Calcote, 2003), while warm and humid summers facilitate its expansion (Graumlich et al., 1989; Whitlock and Bartlein, 1997). This study further confirmed that winter temperatures are conducive to Tsuga survival and that summer and annual temperatures also impact Tsuga distribution.

### 4.2. The extinction of Tsuga in Europe

This study conducted simulations of the current distribution of *Tsuga* in EA, ENA, and WNA, revealing potential regions of *Tsuga* in Europe. Wagner et al.(2017) and Fanal et al.(2021) suggested that *Tsuga*, particularly *T. heterophylla*, has been extensively planted in Europe as an alien species. This indicates that the current climate in Europe may be suitable for the survival of *Tsuga*. The disappearance of *Tsuga* in Europe may have been related to summer precipitation and winter temperature declines during the extremely cold period (440 ka and the LGM).

The potential distribution area of *Tsuga* in the past is aligned with previous palynology research records (Bertini, 2010; Corrado and Magri, 2011; Santangelo et al., 2012; Magri et al., 2017). This alignment suggests that the climate adaptation mechanism of Tsuga in Europe may have been similar to that in EA. Arid climatic conditions affected the survival and distribution of Tsuga, leading to its disappearance in central Italy in the early Brunon period (700 ka, Bertini, 2000). Summer drought in particular has had a significant impact on the survival of Tsuga in Italy (Ravazzi et al., 2005). In addition, the extinction of Tsuga in Europe was related to continuous cooling in the early Miocene and rapid and severe cooling during the Pleistocene, forming a fragmented distribution in EA and NA (Farjon and Filer, 2013). The fossil evidence, particularly the Oligocene fossils T. moenana Kirchheimer and T. plicata (Geinitz) Mai reported in Germany, exhibits strong morphological similarities to the extant Chinese Tsuga species T. chinensis and T. dumosa (Mai and Walther, 1991, 1978). This morphological similarity suggests that Tsuga in Europe may have similar climatic requirements to Tsuga in EA. Therefore, summer precipitation and winter temperature decline during the extremely cold periods (440 ka and the LGM) may have led to Tsuga disappearance in Europe.

This study utilized species distribution modeling, and pollen data provided important results regarding the factors contributing to the disappearance of *Tsuga*. Future research should incorporate fossil morphology data, molecular phylogeny, and physiological experiments to further explore the factors contributing to its disappearance from Europe during the Quaternary.

# 4.3. Distribution changes of potential suitable Tsuga regions during the past

The results of this study indicated that most of the distribution region of Tsuga expanded to lower elevations and latitudes during warm to cold periods due to decreased winter temperatures (Fig. S13a1-d1, a3-d4). Similar cold expansion patterns have also been found in other taxa of Pinaceae. Pinus armandii Franch. (Huang et al., 2023), temperate coniferous forests, are one example (Dakhil et al., 2019), Picea likiangensis (Franch) Pritz, P. purpurea Mast. and P. wilsonii Mast. (Zhang et al., 2018). Climate deterioration led to an increase in the percentage of Tsuga pollen during the last glacial period (28.3-17.6 ka, Chen et al., 2017). The contracting regions have mainly been located in high latitude and elevation regions, characterized by reduced summer precipitation and extreme decreases in winter temperature (Fig. S13a1-d1, a3-d4). Tsuga pollen existed in the Korean Peninsula until the middle Pleistocene, and none existed in the Holocene (Kong, 2000). The disappearance of *Tsuga* in the Korean Peninsula may caused by decreases in winter temperatures (Fig. S13c1, c3-c4). Tsuyama et al. (2014) suggested that T. seiboldii is distributed in temperate regions with warm climates, while T. diversifolia is distributed in colder winter temperature regions, and the increase in summer drought before or during the LGM in the Quaternary glacial periods caused the extinction of T. diversifolia on Hokkaido Island. That is why the Tsuga distribution pattern differs between northern and southern Japan, and the reduction in summer precipitation caused Tsuga to disappear in part of Japan. During the transition from a relatively cold climate to a warm climate, the potentially suitable area of Tsuga in the lower latitudes also contracted (Fig. S13a2-d2, a5-d6). This study found that Tsuga in EA had different response dynamics to climate change during glacial and interglacial periods at different latitudes.

In ENA, the distribution area of Tsuga shifted to lower latitudes and elevations from warm to cold periods due to precipitation and temperature decreases (Fig. S14a1-d1, a3-d4). This finding was in agreement with the observations of Prentice et al. (1991) and Williams (2009) who proposed that the potential distribution of Tsuga was limited to the southern Appalachian Mountains in the LGM. Conversely, during the transition from a relatively cold period to a warm period, accompanied by increases in precipitation and temperature, Tsuga expanded towards high latitudes (Fig. S14a2-d2, a5-d6). This phenomenon has also been confirmed by pollen research; for example, Davis et al. (1986) demonstrated the rapid expansion of Tsuga in the eastern half of upper Michigan at approximately 6-5.5 ka, Jacobson et al. (1987) described expansion routes to the northeast, north, and west, and Prentice et al. (1991) suggested that Tsuga expanded in response to rising winter and summer temperatures and increasing precipitation. Compared with previous studies, this study represented the distribution pattern of Tsuga and the corresponding relationship with important climate changes, revealed Quaternary climate change caused the migration of Tsuga in the latitudes and elevation distribution pattern.

In WNA, *Tsuga* exhibited a similar movement pattern to ENA, transitioning from warm to cold periods and corresponding with decreased winter precipitation and temperatures to move to higher latitudes (Fig. S15a1–d1, a3–d4). For example, *Tsuga* in WNA contracted towards lower latitudes during the LGM (Williams, 2009). Gedalof and Smith (2001) found that *T. mertensiana* is sensitive to summer temperatures in WNA, while it is susceptible to winter precipitation in southern Alaska. Our results showed that in addition to summer temperatures and winter precipitation, winter temperatures also impacted *Tsuga* distribution pattern in WNA. When the climate changed from cold to warm, *Tsu-ga*expanded and migrated to higher latitudes (Fig. S15a2–d2, a5–d5). The warm climate promoted the northward migration of *Tsuga* during postglacial periods, and humid and cool climate conditions further promoted the expansion of *Tsuga* (Rosenberg et al., 2003). Brown (2000) and Lacourse (2012) also stated that the expansion of *Tsuga* were driven by a humid climate environment. This study revealed that the distribution dynamics of *Tsuga* in the WNA were mainly driven by changes in winter precipitation during glacial and interglacial periods, followed by changes in winter and summer temperatures.

# 4.4. Potential distribution in the future and conservation advice

The results of this study indicated that a projected shift in the potentially suitable area for *Tsuga* towards higher latitudes and elevations in the future, with contraction in EA and NA (Fig. S16–S18). This study corroborate previous findings suggesting that *T. canadensis* is expected to expand towards higher latitudes in the northwest region (Dunckel et al., 2017), and *T. mertensiana* is expected to decrease and migrate to higher elevations (Means, 1990). More importantly, our research predicted the distribution pattern based on the responses of different taxa distributed in different regions, providing valuable guidance for future *Tsuga* conservation efforts. Tsuyama et al. (2014) suggested that *T. diversifolia* extinct in Hokkaido in Japan during a certain ice age in the Pleistocene, but was unable to migrate back to Hokkadio Island when the climate was suitable due to the barrier of lowlands and the Tsugaru Straight.

Based on the findings presented above, it is imperative to implement relevant actions to protect *Tsuga* populations in their natural habitats. The study proposes the following strategies:

1) Focused monitoring and reserve adjustments: Give special attention on monitoring the natural growth status of *Tsuga* populations in southwestern China, around the Himalaya, and southern Japan in EA; the southwest Appalachian Mountains in ENA; the Coast Mountains and Rocky Mountains in WNA; and adjust nature reserves or protected regions to better protect *Tsuga* populations.

2) Customized conservation approaches: In areas where the potential distribution of *Tsuga* is threatened, consider implementing *ex-situ* conservation measures. In regions where *Tsuga* populations are expected to increase, explore options for artificial introduction and cultivation. For stable regions, prioritize an *in-situ* conservation strategy, designating these areas as refuges for *Tsuga* and intensifying protective efforts.

3) Prioritizing preservation and mitigation: It is crucial to prioritize assessing the impacts of precipitation changes on the distribution of *Tsuga*. This may contribute to decisions on conservation strategies and help to mitigate potential threats to *Tsuga* populations. Implementing these strategies could contribute to the preservation and safeguarding of *Tsuga* populations, ensuring their long-term survival under changing environmental conditions.

# 5. Conclusions

This study employed climate data and modern distribution data for *Tsuga* to assess the present, past, and future potential distribution regions of the species in EA, ENA, and WNA. The findings highlighted the significant influence of precipitation as well as temperature on the survival and distribution of *Tsuga*. More importantly, our research shows that the seasonal precipitation demand of *Tsuga* varies in different regions. In EA, summer precipitation and winter temperatures exhibited a strong impact on *Tsuga* distribution, while winter precipitation, annual precipitation and winter temperatures played key roles in NA, ENA, and WNA, respectively. Our result indicated that the decreases in summer precipitation and winter temperatures may have led to the extinction of *Tsuga* in Europe. In EA, the potentially suitable area for *Tsuga* was larger during relatively cold periods. In NA, the potentially suitable area

scenarios, the overall distribution area of Tsuga is projected to significantly decrease in EA and NA. These findings emphasize the importance of considering precipitation changes in future conservation efforts for Tsuga populations. It is crucial to vigilantly monitor and protect Tsuga populations in EA and NA. We propose strategies for ex-situ protection, artificial introduction, and in-situ cultivation protection tailored to potential distribution areas with varying changes. Through simulating the historical distribution patterns of Tsuga, our research reveals the possible reasons for the disappearance patterns of Tsuga in Europe and the Korean Peninsula. Furthermore, we suggest that the disappearance time of Tsuga in the Korean Peninsula may occurred in LGM. However, it's worth acknowledging that our study has certain constraints. We concentrated primary on the implications of climate change for species distribution and offer pertinent recommendations. The potential effects of human activities, including alterations in land use, logging, and natural events such as wildfires, could exert effect on the persistence and distribution of Tsuga. Despite these constraints, our results underscore that climate remains the predominant factor influencing the distribution of Tsuga, as evident from our findings. Further research on integrating climate data, distribution models, and conservation strategies is crucial for effectively protecting Tsuga populations and helping them better respond to climate change.

# CRediT authorship contribution statement

Shumei Xiao: Conceptualization, Data curation, Formal analysis, Methodology, Software, Supervision, Writing – original draft, Writing – review & editing. Shufeng Li: Funding acquisition, Project administration, Supervision, Writing – review & editing. Jian Huang: Conceptualization, Writing – review & editing. Xiaojun Wang: Conceptualization, Writing – review & editing. Mengxiao Wu: Conceptualization, Writing – review & editing. Rizwan Karim: Writing – review & editing. Rizwan Karim: Writing – review & editing. Tao Su: Conceptualization, Funding acquisition, Supervision, Writing – review & editing.

# Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

# Data availability

Share data at appendix

# Acknowledgements

We are thankful to all members of the palaeoecology research group, and Dr. Qiuyue Zhang of the biogeography and ecology group. This work was financially supported by the National Natural Science Foundation of China (4231001040), the Young and Middle-aged Academic and Technical Leaders of Yunnan (202305AC160051), and grant from Science and Technology Department of Yunnan Province (YNWR-QNBJ-2019-086), and the 14th Five-Year Plan of Xishuangbanna Tropical Botanical Garden, Chinese Academy of Sciences (E3ZKFF7B, E3ZKFF1K).

# Appendix A. Supplementary data

Supplementary data to this article can be found online at https://doi.org/10.1016/j.ecolind.2023.111533.

#### S. Xiao et al.

### References

Ager, T.A., Matthews, J.V., Yeend, W., 1994. Pliocene terrace gravels of the ancestral Yukon River near Circle, Alaska: Palynology, paleobotany, paleoenvironmental reconstruction and regional correlation. Quat. Int. 22–23, 185–206 https://doi.org/ 10/bzwhqw.

An, Z.S., Clemens, S.C., Ji, S., Qiang, X.K., Jin, Z.D., Sun, Y.B., Prell, W.L., Luo, J.J., Wang, S.M., Xu, H., Cai, Y.J., Zhou, W.J., Liu, X.D., Liu, W.G., Shi, Z.G., Yan, L.B., Xiao, X.Y., Chang, H., Wu, F., Ai, L., Lu, F.Y., 2011. Glacial-Interglacial Indian Summer Monsoon Dynamics. Science 333, 719–723. https://doi.org/10.1126/ science.1203752.

Berntsen, C.M., 1958. Silvical Characteristics of Western Hemlock. Pacific Northwest Forest and Range Experiment Station, US Department of Agriculture, Forest Service.

Bertini, A., 2000. Pollen record from Cole Curti and Cesi: Early and Middle Pleistocene mammal sites in the Umbro–Marchean Apennine Mountains (central Italy). J. Quat. Sci. 15 (8), 825–840. https://doi.org/10.1002/1099-1417(200012)15:8<825::AID-JQS561>3.0.CO;2-6.

Bertini, A., 2010. Pliocene to Pleistocene palynoflora and vegetation in Italy: state of the art. Quat. Int. 225 (1), 5–24. https://doi.org/10.1016/j.quaint.2010.04.025.

Bhattacharyya, A., Dhyani, R., Joshi, R., Shekhar, M., Kuniyal, J.C., Ranhotra, P.S., Singh, S.P., 2023. Is survival of Himalayan Cedar (Cedrus deodara) threatened? An evaluation based on predicted scenarios of its growth trend under future climate change. Sci. Total Environ. 882, 163630 https://doi.org/10.1016/j. scitoteny.2023.163630.

Biltekin, D., Popescu, S.M., Suc, J.P., Quézel, P., Jiménez-Moreno, G., Yavuz, N., Çağatay, M.N., 2015. Anatolia: A long-time plant refuge area documented by pollen records over the last 23million years. Rev. Palaeobot. Palynol. 215, 1–22. https:// doi.org/10.1016/j.revpalbo.2014.12.004.

Brown, K.J., 2000. Late Quaternary vegetation, climate, fire history, and GIS mapping of Holocene climates on southern Vancouver Island, British Columbia, Canada (Doctoral dissertation).

Calcote, R., 2003. Mid-Holocene climate and the hemlock decline: the range limit of Tsuga canadensis in the western Great Lakes region, USA. The Holocene. 13 (2), 215–224 https://doi.org/10/c9cqgp.

Cao, C.C, Su, F.L., Song, F., Yan, H.M., Pang, Q.Z., 2022. Distribution and disturbance dynamics of habitats suitable for Suaeda salsa. Ecol. Indic. 140, 108984. https://doi. org/10.1016/j.ecolind.2022.108984.

Chen, J.X., Liu, Y.G., Shi, X.F., Suk, B.C., Zou, J.J., Yao, Z.Q., 2017. Climate and environmental changes for the past 44 ka clarified by pollen and algae composition in the Ulleung Basin, East Sea (Japan Sea). Quat. Int. 441, 162–173. https://doi.org/ 10.1016/j.quaint.2016.09.052.

Corrado, P., Magri, D., 2011. A late Early Pleistocene pollen record from Fontana Ranuccio (central Italy). J. Quat. Sci. 26 (3), 335–344. https://doi.org/10.1002/ jqs.1459.

Dakhil, M.A., Xiong, Q.L., Farahat, E.A., Zhang, L., Pan, K.W., Pandey, B., Olatunji, O.A., Tariq, A., Wu, X.G., Zhang, A.P., Tan, X., Huang, D., 2019. Past and future climatic indicators for distribution patterns and conservation planning of temperate coniferous forests in southwestern China. Ecol. Indic. 107, 105559 https://doi.org/ 10.1016/j.ecolind.2019.105559.

Davis, M.B., 1983. Quaternary History of Deciduous Forests of Eastern North America and Europe. Ann. Mo. Bot. Gard. 70, 550–563. https://doi.org/10.2307/2992086.

Davis, M.B., Woods, K.D., Webb, S.L., Futyma, R.P., 1986. Dispersal versus climate: Expansion of *Fagus* and *Tsuga* into the Upper Great Lakes region. Vegetatio. 67, 93–103. https://doi.org/10.1007/BF00037360.

Ding, S.T., Wu, J.Y., Tang, D.L., Chen, S.Y., Mo, L.B., Sun, B.N., 2021. Seed cones of Tsuga (Pinaceae) from the upper Miocene of eastern China: Biogeographic and paleoclimatic implications. Rev. Palaeobot. Palynol. 285, 104358 https://doi.org/ 10.1016/j.revpalbo.2020.104358.

Donders, T., Panagiotopoulos, K., Koutsodendris, A., Bertini, A., Mercuri, A.M., Masi, A., Combourieu-Nebout, N., Joannin, S., Kouli, K., Kousis, I., Peyron, O., Torri, P., Florenzano, A., Francke, A., Wagner, B., Sadori, L., 2021. 1.36 million years of Mediterranean forest refugium dynamics in response to glacial-interglacial cycle strength. Proc. Natl. Acad. Sci. U.S.A. 118 (34) https://doi.org/10.1073/ pnas.2026111118 e2026111118.

Dunckel, K., Weiskittel, A., Fiske, G., 2017. Projected Future Distribution of *Tsuga* canadensis across Alternative Climate Scenarios in Maine, U.S. Forests 8 (8), 285. https://doi.org/10.3390/f8080285.

Eiserhardt, W.L., Borchsenius, F., Plum, C.M., Ordonez, A., Svenning, J.C., 2015. Climate-driven extinctions shape the phylogenetic structure of temperate tree floras. Ecol. Lett. 18 (3), 263–272. https://doi.org/10.1111/ele.12409.

Elith, J., Phillips, S.J., Hastie, T., Dudík, M., Chee, Y.E., Yates, C.J., 2011. A statistical explanation of MaxEnt for ecologists. Divers. Distrib. 17 (1), 43–57. https://doi.org/ 10.1111/j.1472-4642.2010.00725.x.

Fanal, A., Mahy, G., Fayolle, A., Monty, A., 2021. Arboreta reveal the invasive potential of several conifer species in the temperate forests of western Europe. NeoBiota. 64, 23–42. https://doi.org/10.3897/neobiota.64.56027.

Farjon, A., Filer, D., 2013. An atlas of the world's conifers: an analysis of their distribution, biogeography, diversity, and conservation status. Brill.

Farjon, A. 2013. Tsuga canadensis. The IUCN Red List of Threatened Species 2013: e. T42431A2979676. doi: 10.2305/IUCN.UK.2013-1.RLTS.T42431A2979676.en. Accessed on 09 August 2023.

Farjon, A. 2013a. Tsuga canadensis. The IUCN Red List of Threatened Species 2013: e. T42431A2979676. https://dx.doi.org/10.2305/IUCN.UK.2013-1.RLTS. T42431A2979676.en. Accessed on 09 August 2023. Farjon, A. 2013b. Tsuga caroliniana. The IUCN Red List of Threatened Species 2013: e. T34200A2850654. https://dx.doi.org/10.2305/IUCN.UK.2013-1.RLTS. T34200A2850654.en. Accessed on 09 August 2023.

Farjon, A., 2017. A handbook of the world's conifers, Second, revised edition. ed. Brill, Leiden; Boston.

Feng, Y.Y., Shen, T.T., Shao, C.C., Du, H., Ran, J.H., Wang, X.Q., 2021. Phylotranscriptomics reveals the complex evolutionary and biogeographic history of the genus *Tsuga* with an East Asian-North American disjunct distribution. Mol. Phylogenet. Evol. 157, 107066 https://doi.org/10.1016/j.jmpev.2020.107066.

Fordham, D.A., Brook, B.W., Moritz, C., Nogués-Bravo, D., 2014. Better forecasts of range dynamics using genetic data. Trends Ecol. Evol. 29 (8), 436–443. https://doi.org/ 10.1016/j.tree.2014.05.007.

Fordham, D.A., Akçakaya, H.R., Alroy, J., Saltré, F., Wigley, T.M.L., Brook, B.W., 2016. Predicting and mitigating future biodiversity loss using long-term ecological proxies. Nat. Clim. Change 6 (10), 909–916. https://doi.org/10.1038/nclimate3086.

Fordham, D.A., Jackson, S.T., Brown, S.C., Huntley, B., Brook, B.W., Dahl-Jensen, D., Gilbert, M.T.P., Otto-Bliesner, B.L., Svensson, A., Theodoridis, S., Wilmshurst, J.M., Buettel, J.C., Canteri, E., McDowell, M., Orlando, L., Pilowsky, J., Rahbek, C., Nogues-Bravo, D., 2020. Using paleo-archives to safeguard biodiversity under climate change. Science 369 (6507), eabc5654. https://doi.org/10.1126/science. abc5654.

Gamisch, A., 2019. Oscillayers: A dataset for the study of climatic oscillations over Plio-Pleistocene time-scales at high spatial-temporal resolution. Global Ecol. Biogeogr. 28 (11), 1552–1560. https://doi.org/10.1111/geb.12979.

GBIF.org (23 August 2022) GBIF Occurrence Download. https://doi.org/10.15468/dl. f2egf9.

Gedalof, Z., Smith, D.J., 2001. Dendroclimatic response of mountain hemlock (Tsuga mertensiana) in Pacific North America. Can. J. for. Res. 31 (2), 322–332. https://doi. org/10.1139/x00-169.

Godman, R., Lancaster, K., 1990. Tsuga canadensis (L.) Carr. eastern hemlock. Silvics of North America 1, 604–612.

Graumlich, L.J., Brubaker, L.B., Grier, C.C., 1989. Long-Term Trends in Forest Net Primary Productivity: Cascade Mountains, Washington. Ecology 70, 405–410. https://doi.org/10.2307/1937545.

Guo, Y.L., Zhao, Z.F., Qiao, H.J., Wang, R., Wei, H.Y., Wang, L.K., Gu, W., Li, X., 2020. Challenges and development trend of species distribution model. Advances in Earth Science 35 (12), 1292–1305. https://doi.org/10.11867/j.issn.1001-8166.2020.110. Hansel, A.K., McKay III, E.D., 2010. Ouaternary period. Geology of Illinois 216–247.

Hayashi, R., 2010. Millennial-scale vegetation changes during the last 40,000yr based on a pollen record from Lake Biwa. Japan. Quat. Res. 74 (1), 91–99. https://doi.org/ 10.1016/j.ygres.2010.04.008.

Hewitt, G., 2000. The genetic legacy of the Quaternary ice ages. Nature 405, 907–913. https://doi.org/10.1038/35016000.

Hewitt, G., 2003. Ice ages: species distributions, and evolution. Evolution on Planet Earth. Academic Press 339–361. https://doi.org/10.1016/B978-012598655-7/ 50045-8.

Holman, G., Del Tredici, P., Havill, N., Lee, N.S., Cronn, R., Cushman, K., Mathews, S., Raubeson, L., Campbell, C.S., 2017. A New Species and Introgression in Eastern Asian Hemlocks (Pinaceae: Tsuga). Syst. Bot. 42, 733–746. https://doi.org/10.1600/ 036364417X696474.

Hong, H.L., Bae, C.J., Zhang, Z.S., 2018. Cenozoic climate change in eastern Asia: Part I. Palaeogeogr. Palaeoclimatol. Palaeoecol. 510, 1–5. https://doi.org/10.1016/j. palaeo. 2018.07.025

Hough, A.F., 1960. Silvical characteristics of eastern hemlock (Tsuga canadensis).132.

Huang, L.L., Li, S.F., Huang, W.Y., Xiang, H.L.L, Jin, J.H., Oskolski, A.A., 2023. Glacial expansion of cold-tolerant species in low latitude: megafossil evidence and distribution modelling. Natl Sci Rev Natl. Sci. Rev. 10 (4), nwad038. https://doi.org/ 10.1093/nsr/nwad038.

Jacobson Jr., G.L., Webb III, T., Grimm, E.C., 1987. Patterns and rates of vegetation change during the deglaciation of eastern North America. North America and adjacent oceans during the last. deglaciation 3, 277–288.

Jacques, J.M., Douglas, M., McAndrews, J.H., 2000. Mid-Holocene hemlock decline and diatom communities in van Nostrand Lake, Ontario, Canada. J Paleolimnol. 23, 385–397.

Katsuki, T., Luscombe, D., 2013. Tsuga sieboldii. The IUCN Red List of Threatened Species 2013: e.T191663A1991616. https://dx.doi.org/10.2305/IUCN.UK.2013-1. RLTS.T191663A1991616.en.

Kong, W.S., 2000. Vegetational history of the Korean Peninsula. Global Ecol. Biogeogr. 9 (5), 391–402. https://doi.org/10.1046/j.1365-2699.2000.00203.x.

Lacourse, T., 2012. A 14,000year vegetation history of a hypermaritime island on the outer Pacific coast of Canada based on fossil pollen, spores and conifer stomata. Quat. Res. 78 (3), 572–582.

Lepage, B.A., 2003. A new species of Tsuga (Pinaceae) from the middle Eocene of Axel Heiberg Island, Canada, and an assessment of the evolution and biogeographical history of the genus. Bot. J. Linn. Soc. 141 (3), 257–296. https://doi.org/10.1046/ j.1095-8339.2003.00131.x.

Li, N., 1995. Studies on the geographic distribution, origin, and dispersal of the family Pinaceae Lindl. Acta PhytotaxonomicaSinica. 33 (02), 105–130.

Li, Y., Gee, C.T., Tan, Z.Z., Zhu, Y.B., Yi, T.M., Li, C.S., 2023. Exceptionally wellpreserved seed cones of a new fossil species of hemlock, Tsuga weichangensis sp. nov. (Pinaceae), from the Lower Miocene of Hebei Province, North China. J. Syst. Evol. https://doi.org/10/gst3cn.

Liu, L., Guan, L.L., Zhao, H.X., Huang, Y., Mou, Q.Y., Liu, K., Chen, T.T., Wang, X.Y., Zhang, Y., Wei, B., Hu, J.Y., 2021. Modeling habitat suitability of Houttuynia cordata Thunb (Ceercao) using MaxEnt under climate change in China. Ecol. Inf. 63, 101324 https://doi.org/10.1016/j.ecoinf.2021.101324. Liu, X.T., Yuan, Q., Ni, J., 2019. Current status of simulation research on plant distribution in China. J. Plant Ecol. 43 (4), 273–283.

- Liu, L., Yuan, Y.F., 2022. Late Quaternary fire and vegetation history inferred from the Xifeng loess-paleosol sequence of the Chinese Loess Plateau. Palaeogeogr. Palaeoclimatol. Palaeoecol. 599, 111072 https://doi.org/10.1016/j. palaeo.2022.111072.
- Magri, D., Di Rita, F., Aranbarri, J., Fletcher, W., González-Sampériz, P., 2017. Quaternary disappearance of tree taxa from Southern Europe: Timing and trends. Quat. Sci. Rev. 163, 23–55. https://doi.org/10.1016/j.quascirev.2017.02.014.
- Magri, D., Palombo, M.R., 2013. Early to Middle Pleistocene dynamics of plant and mammal communities in South West Europe. Quat. Int. 288, 63–72. https://doi.org/ 10.1016/j.quaint.2012.02.028.
- Mai, D.H., Walther, H., 1978. Die Floren der Haselbacher Serie im Weisselster-Becken (Bezirk Leipzig, DDR).
- Mai, D.H., Walther, H., 1991. Die oligozänen und untermiozänen Floren Nordwest-Sachsens und des Bitterfelder Raumes. Abhandlungen Des Staatlichen Museums Für Mineralogie Und Geologie Zu Dresden 38, 1–230.
- McLachlan, J.S., Brubaker, L.B., 1995. Local and regional vegetation change on the northeastern Olympic Peninsula during the Holocene. Can. J. Bot. 73 (10), 1618–1627. https://doi.org/10.1139/b95-175.
- Means, J.E., 1990. Tsuga mertensiana (Bong.) Carr. mountain hemlock. Silvics of North America 1, 623–631.
- Munoz, S.E., Gajewski, K., Peros, M.C., 2010. Synchronous environmental and cultural change in the prehistory of the northeastern United States. Proc. Natl. Acad. Sci. U.S. A. 107 (51), 22008–22013. https://doi.org/10.1073/pnas.1005764107.
- Ni, Z.Y., Jones, R., Zhang, E.L., Chang, J., Shulmeister, J., Sun, W.W., Wang, Y.B., Ning, D.L., 2019. Contrasting effects of winter and summer climate on Holocene montane vegetation belts evolution in southeastern Qinghai-Tibetan Plateau. China. Palaeogeogr. Palaeoclimatol. Palaeoecol. 533, 109232 https://doi.org/10.1016/j. palaeo.2019.06.005.
- Oswald, W.W., Foster, D.R., 2012. Middle-Holocene dynamics of Tsuga canadensis (eastern hemlock) in northern New England, USA. The Holocene. 22 (1), 71–78 https://doi.org/10/brwdgc.
- Phillips, S.J., Anderson, R.P., Schapire, R.E., 2006. Maximum entropy modeling of species geographic distributions. Ecol. Modell. 190 (3–4), 231–259. https://doi.org/ 10.1016/j.ecolmodel.2005.03.026.
- Phillips, S.J., Anderson, R.P., Dudík, M., Schapire, R.E., Blair, M.E., 2017. Opening the black box: an open-source release of Maxent. Ecography 40 (7), 887–893. https:// doi.org/10.1111/ecog.03049.
- Phillips, S.J., Dudík, M., 2008. Modeling of species distributions with Maxent: new extensions and a comprehensive evaluation. Ecography 31 (2), 161–175. https://doi. org/10.1111/j.0906-7590.2008.5203.x.
- Prentice, I.C., Bartlein, P.J., Webb III, T., 1991. Vegetation and Climate Change in Eastern North America Since the Last Glacial Maximum. Ecology 72 (6), 2038–2056. https://doi.org/10.2307/1941558.
- Qiao, C.Y., Ran, J.H., Li, Y., Wang, X.Q., 2007. Phylogeny and Biogeography of Cedrus (Pinaceae) Inferred from Sequences of Seven Paternal Chloroplast and Maternal Mitochondrial DNA Regions. Ann. Bot. 100 (3), 573–580. https://doi.org/10.1093/ aob/mcm134.
- Ravazzi, C., Pini, R., Breda, M., Martinetto, E., Muttoni, G., Chiesa, S., Confortini, F., Egli, R., 2005. The lacustrine deposits of Fornaci di Ranica (late Early Pleistocene, Italian Pre-Alps): stratigraphy, palaeoenvironment and geological evolution. Quat. Int. 131 (1), 35–58. https://doi.org/10.1016/j.quaint.2004.07.021.
- Int. 131 (1), 35–58. https://doi.org/10.1016/j.quaint.2004.07.021.
  Rosenberg, S.M., Walker, I.R., Mathewes, R.W., 2003. Postglacial spread of hemlock (Tsuga) and vegetation history in Mount Revelstoke National Park, British Columbia. Canada. Can. J. Bot. 81 (2), 139–151 https://doi.org/10/dnq22j.
- Sandel, B., Arge, L., Dalsgaard, B., Davies, R.G., Gaston, K.J., Sutherland, W.J., Svenning, J.C., 2011. The influence of Late Quaternary climate-change velocity on species endemism. Science 334 (6056), 660–664. https://doi.org/10.1126/ science.1210173.
- Santangelo, N., Di Donato, V., Lebreton, V., Romano, P., Russo Ermolli, E., 2012. Palaeolandscapes of Southern Apennines during the late Early and the Middle Pleistocene. Quat. Int. 267, 20–29. https://doi.org/10.1016/j.quaint.2011.02.036.
- Sun, A.Z., Feng, Z.D., 2015. Climatic charges in the western part of the Chinese Loess Plateau during the Last Deglacial and the Holocene: a synthesis of pollen records. Quat. Int. 372, 130–141. https://doi.org/10.1016/j.quaint.2014.10.046.
- Svenning, J.C., Fløjgaard, C., Marske, K.A., Nógues-Bravo, D., Normand, S., 2011. Applications of species distribution modeling to paleobiology. Quat. Sci. Rev. 30 (21–22), 2930–2947. https://doi.org/10.1016/j.quascirev.2011.06.012.
- Tarasov, P.E., Nakagawa, T., Demske, D., Österle, H., Igarashi, Y., Kitagawa, J., Mokhova, L., Bazarova, V., Okuda, M., Gotanda, K., Miyoshi, N., Fujiki, T., Takemura, K., Yonenobu, H., Fleck, A., 2011. Progress in the reconstruction of Quaternary climate dynamics in the Northwest Pacific: A new modern analogue reference dataset and its application to the 430-kyr pollen record from Lake Biwa. Earth Sci. Rev. 108 (1–2), 64–79. https://doi.org/10.1016/j.earscirev.2011.06.002.

- Thompson, R., Anderson, K., Bartlein, P., 1999. Atlas of relations between climatic parameters and distributions of important trees and shrubs in North America. US Department of the Interior, US Geological Survey.
- Toney, J.L., Rodbell, D.T., Miller, N.G., 2003. Sedimentologic and palynologic records of the last deglaciation and Holocene from Ballston Lake. New York. Quat. Res. 60 (2), 189–199. https://doi.org/10.1016/S0033-5894(03)00093-0.
- Tsukada, M., 1985. Map of vegetation during the last glacial maximum in Japan. Quat. Res. 23 (3), 369–381. https://doi.org/10.1016/0033-5894(85)90041-9.
- Tsuyama, I., Nakao, K., Higa, M., Matsui, T., Shichi, K., Tanaka, N., 2014. What controls the distribution of the Japanese endemic hemlock, Tsuga diversifolia? Footprint of climate in the glacial period on current habitat occupancy. J. for. Res. 19 (1), 154–165. https://doi.org/10.1007/s10310-013-0399-9.
- Van Andel, T.H., Tzedakis, P.C., 1996. Palaeolithic landscapes of Europe and environs, 150,000–25,000 years ago: an overview. Quat. Sci. Rev. 15 (5–6), 481–500 https:// doi.org/10/b86ffk.
- Varela, S., Lobo, J.M., Hortal, J., 2011. Using species distribution models in paleobiogeography: a matter of data, predictors and concepts. Palaeogeogr. Palaeoclimatol. Palaeoecol. 310 (3–4), 451–463. https://doi.org/10.1016/j. palaeo.2011.07.021.
- Vasiliauskas, S., 1996. Interpretation of age-structure gaps in hemlock (Tsuga canadensis) populations of Algonquin Park. Queen's University at Kingston.
- Wagner, V., Chytrý, M., Jiménez-Alfaro, B., Pergl, J., Hennekens, S., Biurrun, I., Knollová, I., Berg, C., Vassilev, K., Rodwell, J.S., Škvorc, Ž., Jandt, U., Ewald, J., Jansen, F., Tsiripidis, I., Botta-Dukát, Z., Casella, L., Attorre, F., Rašomavičius, V., Ćušterevska, R., Schaminée, J.H.J., Brunet, J., Lenoir, J., Svenning, J.C., Kącki, Z., Petrášová-Šibíková, M., Šilc, U., García-Mijangos, I., Campos, J.A., Fernández-González, F., Wohlgemuth, T., Onyshchenko, V., Pyšek, P., 2017. Alien plant invasions in European woodlands. Divers. Distrib. 23 (9), 969–981. https://doi.org/ 10.1111/ddi.12592.
- Wang, H.B., Bondarenko, O.V., Jacques, F.M.B., Wang, Y.H., Zhou, Z.K., 2015. A New Species of Tsuga (Pinaceae) based on Lignified Wood from the Late Miocene of Central Yunnan, China, and Its Paleoenvironmental Implications. Acta Geologica Sinica - English Edition 89 (5), 1429–1439. https://doi.org/10.1111/1755-6724.12555.
- Whitlock, C., Bartlein, P.J., 1997. Vegetation and climate change in northwest America during the past 125 kyr. Nature. 388 (6637), 57–61. https://doi.org/10.1038/ 40380
- Willeit, M., Ganopolski, A., Calov, R., Brovkin, V., 2019. Mid-Pleistocene transition in glacial cycles explained by declining CO<sub>2</sub> and regolith removal. Sci. Adv. 5 (4), eaav7337. https://doi.org/10.1126/sciadv.aav7337.
- Williams, J.W., Shuman, B.N., Webb III, T., Bartlein, P.J., Leduc, P.L., 2004. Late-Quaternary vegetation dynamics in North America: scaling from taxa to biomes. Ecol. Monogr. 74 (2), 309–334. https://doi.org/10.1890/02-4045.
- Williams, J., 2009. Quaternary Vegetation Distribution. Encyclopedia of Paleoclimatology and Ancient Environments. Dordrecht, The Netherlands: Springer Netherlands. 856–861. doi: 10.1007/978-1-4020-4411-3 199.
- Wu, M.X., Huang, J., Su, T., Leng, Q., Zhou, Z.K., 2020. *Tsuga* seed cones from the late Paleogene of southwestern China and their biogeographical and paleoenvironmental implications. Palaeoworld 29 (3), 617–628. https://doi.org/10.1016/j. palwor.2019.07.005.
- Xiao, S.M., Li, S.F., Wang, X.J., Chen, L.L., Su, T., 2022. Cedrus distribution change: past, present, and future. Ecol. Indic. 142, 109159 https://doi.org/10.1016/j. ecolind.2022.109159.
- Xing, Y.W., Liu, Y.C., Su, T., Jacques, F.M., Zhou, Z.K., 2013. A new *Tsuga* species from the upper Miocene of Yunnan, southwestern China and its palaeogeographic significance. Palaeoworld 22 (3–4), 159–167. https://doi.org/10.1016/j. palwor.2013.09.003.
- Xu, D.K., Lu, H.Y., Wu, N.Q, Liu, Z.X., 2010. 30,000-Year vegetation and climate change around the East China Sea shelf inferred from a high-resolution pollen record. Quat. Int. 227 (1), 53–60. https://doi.org/10.1016/j.quaint.2010.04.015.
- Yang, Y. & Luscombe, D. 2013. Tsuga forrestii. The IUCN Red List of Threatened Species 2013: e.T32457A2819347. doi: 10.2305/IUCN.UK.2013-1.RLTS.T32457A2819347. en. Accessed on 09 August 2023.
- Yang Q.S., Xing Y.W., Zhou Z.K., 2009. Modern distribution and climatic characteristics of *Tsuga* in Asian monsoon region. Acta Botanica Yunnanica. 31(5), 389–398. doi: 10.3724 SP.J.1143.2009.
- Yang, Y., Xu, X.T., 2017. Classification and geographical distribution of gymnosperms worldwide. Shanghai Science and Technology Press.
- Zhang, A.P., Wang, Y., Xiong, Q.L., Wu, X.G., Sun, X.M., Huang, Y.M., Zhang, L., Pan, K. W., 2018. Historical distribution and refuge of three species of Spruce since the last Interglacial. J. Appl. Ecol. 29 (07), 2411–2421. https://doi.org/10.13287/j.1001-9332.201807.027.
- Zhou, Z.K., Huang, J., Ding, W.N., 2017. The impact of major geological events on Chinese flora. Biodivers. Sci. 25, 17–23. https://doi.org/10.17520/biods.2016120.